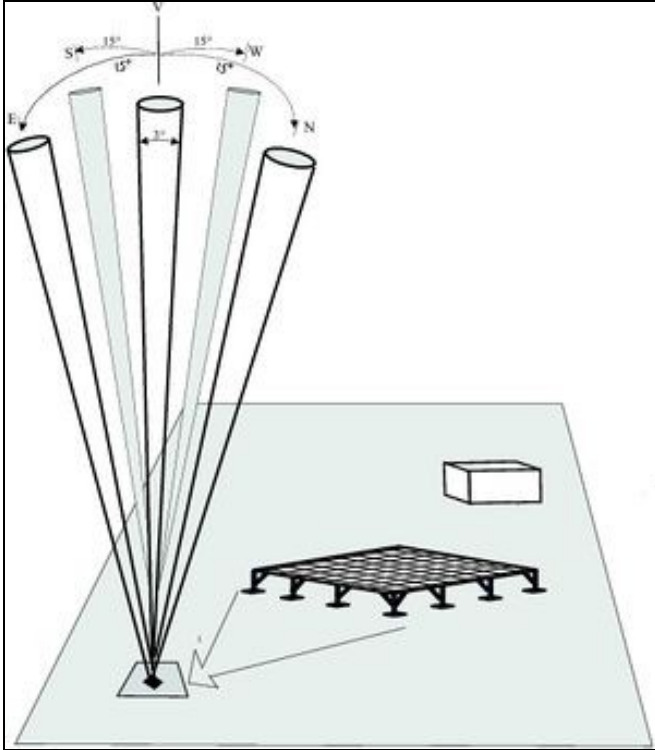


At least three linear independent beam directions and some assumptions concerning the wind field are required to transform the measured 'line-of-sight' radial velocities into the wind vector. Comparisons of RWP winds with data from a meteorological tower ([Adachi et.al., 2005](#)) and balloon soundings ([Rao et.al. 2008](#)) have shown that a four-beam based DBS sampling configuration is superior over a three-beam configuration in terms of data quality.

[Cheong et.al. \(2008\)](#) have found that the RMS error of RWP measurements can be significantly reduced by increasing the number of off-vertical beams in DBS beyond four. At present, such a configuration can only be used by a few RWP's systems because of restrictions imposed by the simple phased array constructions that are mostly used.



### 5 beam DBS profiler antenna pointing

For a five-beam system and a (horizontally) constant wind field, the relation between wind vector and the 'line of sight' radial wind components is given by the following linear system:

$$\begin{pmatrix} \sin(\alpha_1)\sin(\epsilon_1) & \cos(\alpha_1)\sin(\epsilon_1) & \cos(\epsilon_1) \\ \sin(\alpha_2)\sin(\epsilon_2) & \cos(\alpha_2)\sin(\epsilon_2) & \cos(\epsilon_2) \\ \sin(\alpha_3)\sin(\epsilon_3) & \cos(\alpha_3)\sin(\epsilon_3) & \cos(\epsilon_3) \\ \sin(\alpha_4)\sin(\epsilon_4) & \cos(\alpha_4)\sin(\epsilon_4) & \cos(\epsilon_4) \\ \sin(\alpha_5)\sin(\epsilon_5) & \cos(\alpha_5)\sin(\epsilon_5) & \cos(\epsilon_5) \end{pmatrix} \begin{pmatrix} u \\ v \\ w \end{pmatrix} = \begin{pmatrix} v_{r1} \\ v_{r2} \\ v_{r3} \\ v_{r4} \\ v_{r5} \end{pmatrix}$$

where  $\alpha_i$  and  $\epsilon_i$  denote azimuth and elevation angles of beam  $i$ . In compact matrix notation, this can be written as

$$A\vec{v} = \vec{v}_r.$$

This over-determined system can be solved in a least-squares sense and the wind vector components can be obtained from the measured radial velocities through a pseudo-inverse as (the matrix superscripts T and I denote transposition and inverse, respectively).:

$$\vec{v} = (A^T A)^{-1} A^T \vec{v}_r$$

The homogeneity condition is of course not always fulfilled, in particular not in a convective boundary layer or in patchy precipitation. A discussion of the DBS method can be found in [Koscielny et.al. \(1984\)](#), [Weber et.al. \(1992\)](#), [Goodrich et. al. \(2002\)](#). The problem and the resulting measurement errors have recently been investigated by [Scipion et.al. \(2009\)](#), they are even noticeable in NWP data assimilation ([Cardinali, 2009](#)). However, the assumption is usually assumed to be correct for winds averaged over a longer time interval. [Cheong et.al. \(2008\)](#) have used data obtained with a volume-imaging multi-signal wind profiler in a convective boundary layer to show that for this particular case the assumptions inherent in the DBS method were valid for a wind field averaged over 10~minutes. This is the main reason why DBS RWP wind measurements are typically averages over 10-60~minutes. More work is required to obtain reliable quantitative estimates of this error under a variety of atmospheric conditions.